

## Nonlinear stability analysis in multilayer quasigeostrophic systems

This article has been downloaded from IOPscience. Please scroll down to see the full text article.

1981 J. Phys. A: Math. Gen. 14 L203

(<http://iopscience.iop.org/0305-4470/14/5/019>)

View [the table of contents for this issue](#), or go to the [journal homepage](#) for more

Download details:

IP Address: 129.252.86.83

The article was downloaded on 31/05/2010 at 06:09

Please note that [terms and conditions apply](#).

**LETTER TO THE EDITOR**

**Nonlinear stability analysis in multilayer quasigeostrophic systems**

Stefano Pierini† and Angelo Vulpiani

Istituto di Fisica G Marconi, University of Rome, Italy

Received 2 March 1981

**Abstract.** A sufficient condition for nonlinear stability of steady solutions of the quasigeostrophic equations in a multilayer system is found. It proves the stability of given shear flows in a two-layer model that were known to be neutrally stable only.

In this letter we study the problem of nonlinear stability of steady-state solutions in a multilayer quasigeostrophic fluid dynamical system. By a generalisation of a theorem (Arnol'd 1965) about nonlinear stability of stationary plane curvilinear flows, we obtain a sufficient condition for a steady solution to be stable in the Lyapunov sense, i.e. without linearising the evolution equations in the perturbation. The application of our result to a two-layer shear flow gives a rigorous stability condition which was previously obtained through linear computations.

The large scale motions of oceans and atmospheres are described (Pedlosky 1979) by the shallow-water equations in the quasigeostrophic approximation (small Rossby number). If the stratification in density is modelled by  $N$  superimposed layers whose constant densities are  $\rho_1 < \rho_2 < \dots < \rho_N$  (the first layer is the upper one) the evolution is given by the following system of partial differential equations (Pedlosky 1979)

$$\frac{\partial q_i}{\partial t} + J(\psi_i, q_i) = 0 \quad i = 1, \dots, N$$

$$J(a, b) = a_x b_y - a_y b_x \quad + \text{boundary and initial conditions} \quad (1)$$

where  $\psi_i$  is the stream function of the  $i$ th layer (i.e. the horizontal velocity field in the  $i$ th layer is  $v_i(x, y, t) = (-\partial\psi_i/\partial y, \partial\psi_i/\partial x)$ ) and

$$q_i = \Delta\psi_i + F_i \sum_{j=1}^N T_{ij}\psi_j + f_i \quad F_i = \frac{f_0^2}{g[(\rho_{i+1} - \rho_i)/\rho_0]D_i}$$

$$T_{ij} = \begin{pmatrix} -1 & 1 & & & 0 \\ 1 & -2 & 1 & & \\ & & & \dots & \\ 0 & & & 1 & -2 & 1 \\ & & & & 1 & -1 \end{pmatrix} \quad f_i = f_0 + \beta y + \delta_{iN} b(y)$$

$$f_0 = 2\Omega \sin \varphi_0 \quad \beta = 2\Omega \cos \varphi_0 / R \quad b = f_0 d(y) / D_N$$

† Present address: Department of Applied Mathematics and Theoretical Physics, University of Cambridge, England.

where  $g$  is the acceleration of gravity,  $\rho_0$  the mean density,  $D_i$  the mean thickness of the  $i$ th layer,  $R$  the Earth's radius,  $\Omega$  the Earth's angular velocity,  $\varphi_0$  the reference latitude and  $d(y)$  the bottom topography.

The steady solutions of equations (1) are given by

$$J(\bar{\psi}_i, \bar{q}_i) = 0 \quad i = 1, \dots, N \quad + \text{boundary conditions.} \quad (2)$$

A solution of equations (2) is given by the equations

$$\bar{q}_i = P_i(\bar{\psi}_i) \quad i = 1, \dots, N \quad (3)$$

where  $P_i$  is any smooth function such that the boundary conditions can be satisfied. In a closed domain  $\mathcal{C}$  the natural boundary conditions are

$$\psi_i|_{\partial\mathcal{C}} = c_i = \text{constant.}$$

However, in geophysical situations the open boundary conditions are imposed:

$$\lim_{(x,y) \rightarrow \pm\infty} \nabla\psi_i = 0. \quad (4)$$

Usually the study of stability of a steady solution of the differential problem

$$\frac{\partial f}{\partial t} = Af \quad + \text{boundary and initial conditions} \quad (5)$$

(where  $A$  is a nonlinear operator) is carried out in the framework of the linearised stability theory (Joseph 1976). From a given steady solution  $\bar{f}$  ( $A\bar{f} = 0$ ) one can define the linear operator  $\tilde{A}_{\bar{f}}$  by

$$A(\bar{f} + \delta f) = \tilde{A}_{\bar{f}}\delta f + O(\delta f^2); \quad (6)$$

then the usual linearised stability theory consists in analysing the eigenvalues of the linear operator  $\tilde{A}_{\bar{f}}$ . If all the eigenvalues of  $\tilde{A}_{\bar{f}}$  have a negative real part, the steady solution  $\bar{f}$  of equation (5) is asymptotically stable. If one (or more than one) of them has a positive real part,  $\bar{f}$  is unstable. If some eigenvalue is a pure imaginary number and the other eigenvalues are real negatives (neutral stability), no definitive conclusion can be obtained. This last case is, unfortunately, typical in inviscid fluid dynamical problems (see Joseph (1976) for a discussion of this point). Therefore, in this case, nonlinear techniques should be used to discuss stability.

In this context very few results are known and many of them are obtained by applying Lyapunov's well known second theorem. Let  $\mathcal{T}$  be a metric space in which a time evolution  $T^t$  is given.  $\bar{f} \in \mathcal{T}$  is a rest point if  $T^t\bar{f} = \bar{f}$ ,  $\forall t$ . The above mentioned Lyapunov theorem states (Hahn 1963) the following.

*Theorem.* If a functional  $H[f]$  exists such that (a)  $H$  has a relative minimum in  $\bar{f}$  and (b)

$$\frac{d}{dt}H[T^t f] \leq 0 \quad \text{if } f \in I_{\bar{f}}$$

where  $I_{\bar{f}}$  is a neighbourhood of  $\bar{f}$ , then  $\bar{f}$  is a stable rest point in the a Lyapunov sense. In this case  $H$  is called a Lyapunov functional. In our case a point  $f$  corresponds to a vector  $\psi \equiv (\psi_1, \dots, \psi_N)$  and the time evolution  $T^t$  is determined by equation (1).

Let us emphasise that the explicit determination of a Lyapunov functional for a given system is very difficult and it is possible only in some cases in which there are

particular symmetries or conservation laws. In our case, as an infinite number of conservation laws exists, it is possible to determine a Lyapunov functional for some steady solution of the problem.

Let us now present our main result: one can affirm that the steady solution  $\bar{\psi} = (\bar{\psi}_1, \dots, \bar{\psi}_N)$  satisfying equations (3) and (4) with given functions  $P_i$  is stable in the Lyapunov sense if the following conditions hold:

$$P'_i(\bar{\psi}_i) > 0 \quad i = 1, \dots, N. \tag{7}$$

This sufficient condition for stability can be obtained as follows: defining

$$\begin{aligned} \Phi'_i(z) &= P_i^{-1}(z) \quad \text{i.e. } \Phi'_i(P_i(z)) = z \\ H[\psi] &= \int \left( \sum_{i=1}^N \frac{1}{F_i} [\frac{1}{2}(\nabla\psi_i)^2 + \Phi_i(q_i)] + \sum_{i=1}^{N-1} \frac{1}{2}(\psi_i - \psi_{i+1})^2 \right) dx dy \end{aligned}$$

we want to prove that  $H$  is a Lyapunov functional for the steady solution  $\bar{\psi}$  under conditions (7); then, for the second Lyapunov theorem,  $\bar{\psi}$  is a stable steady solution. We thus have to prove that conditions (a) and (b) hold.

(a) The functional  $H$  has a minimum in  $\bar{\psi}$  because

$$\delta H[\bar{\psi}] = 0 \quad \delta^2 H[\bar{\psi}] > 0$$

where  $\delta H$  and  $\delta^2 H$  denote first and second variation of  $H$ . Indeed one has

$$\begin{aligned} \delta H[\psi] &= \int \left( \sum_{i=1}^N \frac{1}{F_i} (\nabla\psi_i \nabla\delta\psi_i + \Phi'_i(q_i)\delta q_i) + \sum_{i=1}^{N-1} (\psi_i - \psi_{i+1})(\delta\psi_i - \delta\psi_{i+1}) \right) dx dy \\ &= \int \left( \sum_{i=1}^N \frac{1}{F_i} \Delta\delta\psi_i (\Phi'_i(q_i) - \psi_i) \right. \\ &\quad \left. + \sum_{i=1}^{N-1} (\delta\psi_i - \delta\psi_{i+1})(\psi_i - \Phi'_i(q_i) - \psi_{i+1} + \Phi'_{i+1}(q_{i+1})) \right) dx dy \end{aligned}$$

$$\delta^2 H[\psi] = \frac{1}{2} \int \left( \sum_{i=1}^N \frac{1}{F_i} [(\nabla\delta\psi_i)^2 + \Phi''_i(q_i)\delta q_i^2] + \sum_{i=1}^{N-1} (\delta\psi_i - \delta\psi_{i+1})^2 \right) dx dy.$$

The first variation is zero for  $\psi = \bar{\psi}$  since  $\bar{\psi}_i = \Phi'_i(\bar{q}_i)$ ; moreover  $\delta^2 H$  is positive for  $\psi = \bar{\psi}$  because of condition (7), which can also be written as

$$\Phi''_i(\bar{q}_i) > 0.$$

(b) The second condition of the theorem is verified because  $H[\psi]$  is conserved in time, i.e.

$$\frac{d}{dt} H[\psi] = 0.$$

The functionals  $\int \Phi_i(q_i) dx dy$  are time invariant for any  $\Phi(\cdot)$  because of the conservation of the quantity  $q_i$  related to a fluid particle in the  $i$ th layer, expressed by (1). Besides, the total energy

$$E = \frac{1}{2} \int \left( \sum_{i=1}^N \frac{1}{F_i} (\nabla\psi_i)^2 + \sum_{i=1}^{N-1} (\psi_i - \psi_{i+1})^2 \right) dx dy$$

is also conserved. This can be seen by multiplying the  $i$ th equation in (1) by  $\psi_i/F_i$  and integrating:

$$\begin{aligned} 0 &= \sum_{i=1}^N \frac{1}{F_i} \int \left( \psi_i \frac{\partial q_i}{\partial t} + \psi_i J(\psi_i, q_i) \right) dx dy \\ &= \sum_{i,j=1}^N \int \frac{1}{F_i} \psi_i \frac{\partial}{\partial t} (\Delta \psi_i + F_i T_{ij} \psi_j) dx dy \\ &= -\frac{1}{2} \frac{d}{dt} \int \left( \sum_{i=1}^N \frac{1}{F_i} (\nabla \psi_i)^2 + \sum_{i=1}^{N-1} (\psi_i - \psi_{i+1})^2 \right) dx dy. \end{aligned}$$

The invariance of  $H$  is thus proved, so  $\bar{\psi}$  is stable if conditions (7) are verified.

As an application let us consider a shear flow described in a two-layer system by the steady solution

$$\bar{\psi}_i = U_i y \quad i = 1, 2 \quad U_i > 0. \quad (8)$$

This solution does not satisfy the boundary condition (4). However, the unmodified conditions for stability (7) hold also in this case providing the domain is a cylinder with its axis parallel to the earth's rotation axis; the damping of the velocity at  $y \rightarrow \pm\infty$  must also be prescribed.

$$\begin{aligned} P'_1(\bar{\psi}_1) &= \frac{\nabla(\Delta \bar{\psi}_1 + F_1(\bar{\psi}_2 - \bar{\psi}_1) + \beta y)}{\nabla \bar{\psi}_1} > 0 \\ P'_2(\bar{\psi}_2) &= \frac{\nabla(\Delta \bar{\psi}_2 + F_2(\bar{\psi}_1 - \bar{\psi}_2) + \beta y + b)}{\nabla \bar{\psi}_2} > 0. \end{aligned} \quad (9)$$

Then if the velocities satisfy

$$-\frac{\beta}{F_1} < U_2 - U_1 < \frac{B}{F_2} \quad B = \beta + b' \quad (10)$$

stability is ensured rigorously, i.e. without having linearised the equations of motion.

If  $U_1, U_2 < 0$ , i.e. if the velocities are directed eastward, conditions (9) can be applied in a moving reference frame. Let us make a Galilean transformation

$$x = x' + ct \quad y = y' \quad t = t'$$

with  $c > \max(|U_1|, |U_2|)$  and  $c \ll R/T$  where  $T$  is the Earth's period of rotation; equations (1) become

$$\frac{\partial \tilde{q}_i}{\partial t'} + J'(\tilde{\psi}_i, \tilde{q}_i) = 0$$

where

$$\tilde{\psi}_i = \psi'_i + cy' \quad \tilde{q}_i = \Delta \tilde{\psi}_i + F_i \sum_{j=1}^N T_{ij} \tilde{\psi}_j + \beta' y' + \delta_{i2} b$$

with  $\beta' \approx \beta$  because  $c \ll R/T$ .

In this new frame our steady solution is

$$\tilde{\psi}_i = (c - U_i)y' = U'_i y' \quad U'_i > 0$$

and so our stability condition can be applied. One obtains

$$-\beta/F_1 < U_1 - U_2 < \beta/F_2. \quad (11)$$

Conditions (10) and (11) have an immediate physical meaning.

In order to study the stability properties of the steady solution (8), Pedlosky (1964), Gill *et al* (1974) and Tang (1979) linearised the evolution equations in the perturbation  $\delta\psi$  and obtained that, under the same condition (11),  $\delta\psi(x, y, t)$  remains bounded but does not decrease in time; thus stability was not inferred rigorously, as discussed at the beginning of this letter. Here, on the contrary, condition (11) (and (10) for  $U_i > 0$ ) is a sufficient condition for rigorous stability (in the Lyapunov sense) of  $\bar{\psi}_i$ .

Thanks are due to Dr P Henrotay for useful discussions and to Drs L Peliti and E Salusti for their criticism of the manuscript.

### References

- Arnol'd V I 1965 *Sov. Math. Dokl.* **6** 773  
Gill A E, Green J S A and Simmons A J 1974 *Deep-Sea Res.* **21** 499  
Hahn W 1963 *Theory and Application of Lyapunov's Direct Method* Englewood Cliffs, NJ: Prentice-Hall)  
Joseph D D 1976 *Stability of Fluid Motion* (Berlin: Springer)  
Pedlosky J 1964 *J. Atmos. Sci.* **21** 201  
——— 1979 *Geophysical Fluid Dynamics* (Berlin: Springer)  
Tang C L 1979 *J. Fluid Mech.* **93** 379